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Using the 2011 M=9.0 Tohoku earthquake to test the Coulomb stress triggering hypothesis and to calculate faults brought closer to failure

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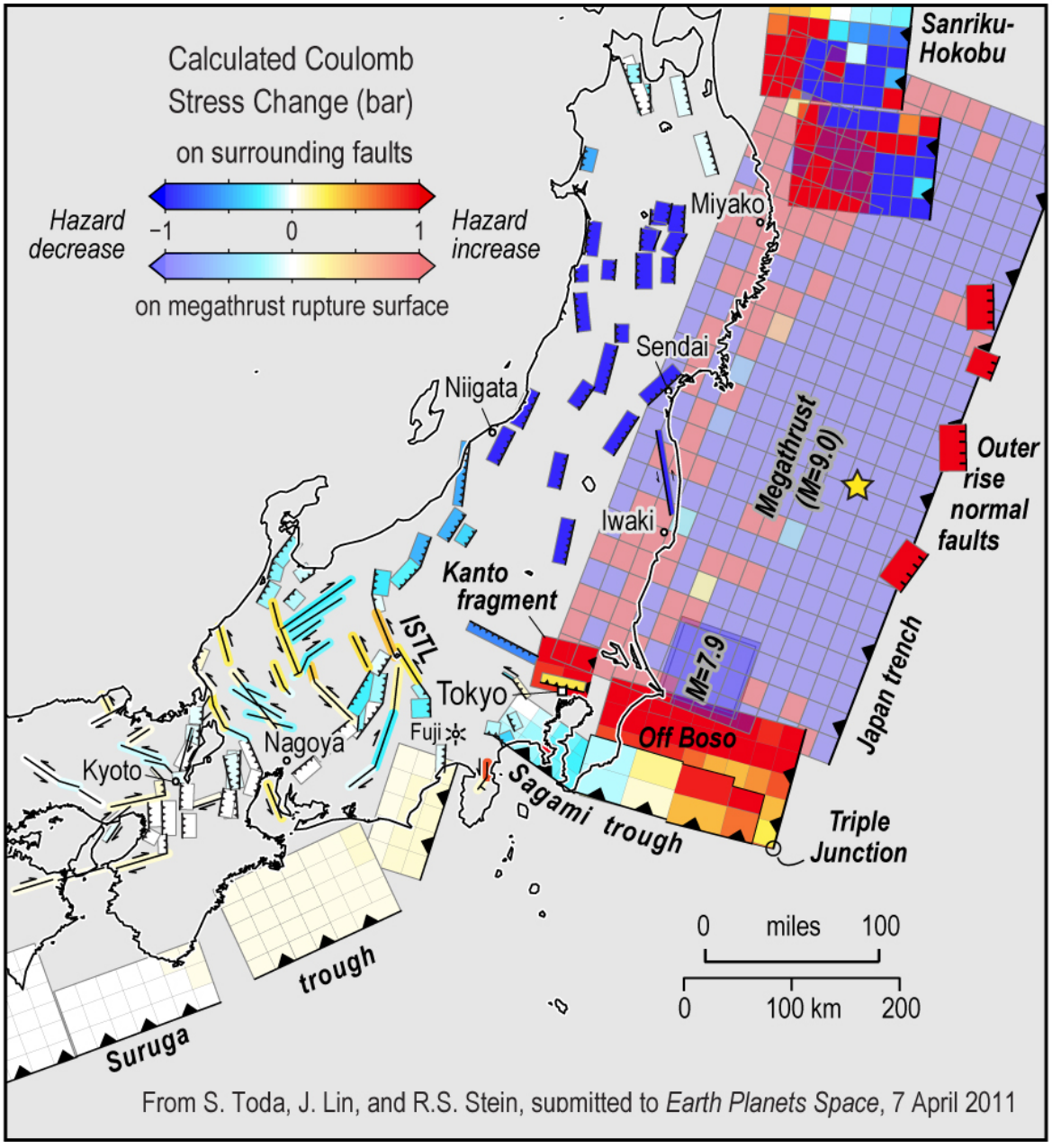
Plain English Summary

Most seismologists assume that once a major earthquake and its expected aftershocks do their damage, the fault will remain quiet until stresses in Earth's crust have time to rebuild, typically over hundreds or thousands of years. But research over the past two decades has shown that earthquakes interact in ways never before imagined. A major shock does relieves stress—and thus the likelihood of a second major tremor—but only in some areas. The probability of a succeeding earthquake adjacent to the section of the fault that ruptured or on a nearby but different fault can jump by as much as a factor of ten.

At the heart of this hypothesis—known as Coulomb stress triggering—is the realization that faults are responsive to subtle stresses they acquire as neighboring faults shift, deforming the Earth's crust. Drawing on records of past tremors and novel calculations of fault behavior, stress relieved during an earthquake does not simply dissipate; instead it moves down the fault and concentrates in sites nearby, promoting subsequent tremors. Indeed, studies of about two dozen faults since 1992 have convinced many of us that aftershocks can be triggered even when the stress is increased by as little as one-tenth the

pressure required to inflate a car tire, or about 0.3 bar. While the frequency of aftershocks decays rapidly with time after the mainshock, their magnitude does not decay at all, and so large late aftershocks are an unexpected consequence of a great earthquake.



To calculate the Coulomb stress transfer, one needs to know the distribution of slip on the mainshock rupture, and the geometry, friction, and sense of slip on surrounding faults. The 11 March 2011 Tohoku earthquake provides an unprecedented test of the extent to which Coulomb stress transfer governs the triggering of aftershocks, and by extension,

subsequent mainshocks on surrounding faults. Thanks to Japan's superb monitoring networks, the earthquake is the best-recorded great event the world has ever known. During 11-31 March 2011, there were almost 200 aftershocks for which the slip and geometry on the causative fault is known (they possess "focal mechanisms"), and so the Coulomb stress change imparted by the mainshock rupture can be resolved on the aftershock fault planes to test whether they were, in fact, brought closer to failure, consistent with the theory.

The distribution of fault slip in the mainshock is nevertheless uncertain; many "source models" for the mainshock have been developed during the past 3 weeks from seismic, geodetic, and tsunami observations. Here we show that among the six source models we tested, there is a mean 47% gain in positively stressed aftershock mechanisms over that for the background earthquakes (shocks that struck before the Tohoku mainshock), which serve as the control group. This is a strong confirmation of the Coulomb hypothesis. Among all the tested models, that of *Wei and Sladen* [2011] not only produced the largest gain, 63%, but the largest yet found for any earthquake we or others have tested. Further, a value of friction (0.4), midway between 'teflon and rough cohesive faults, is found to fit the aftershock set best.

With this test as confirmation, we use the best-performing model and best value of friction to calculate the stress transferred to all major faults that surround the mainshock. We find that large sections of the Japan trench megathrust to the north of the mainshock (the offshore 'Sanriku-Hokobu' region 200 km northeast of Sendai), and to the south of the mainshock (the 'Off-Boso' region 100 km east of Tokyo), were also brought $\geq$0.3 bars closer to failure. East of the megathrust, tensional ('outer rise normal') faults are bought 1-15 bars closer to failure; large earthquakes on these could trigger a tsunami. The 'Kanto fragment' that we believe is wedged 50-70 km beneath Tokyo, and the Itoigawa-Shizuoka Tectonic Line ('ISTL') 100 km west of Tokyo, were also brought $\geq$0.3 bars closer to failure. Based on our other studies, these stress increases are large enough to increase the likelihood of triggering significant aftershocks or subsequent mainshocks.

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Abstract

The 11 March 2011 Tohoku earthquake provides an unprecedented test of the extent to which Coulomb stress transfer governs the triggering of aftershocks. During 11-31 March, there were 177 aftershocks with focal mechanisms, and so the Coulomb stress change imparted by the rupture can be resolved on the aftershock nodal planes to learn whether they were brought closer to failure. Numerous source models for the mainshock have been inverted from seismic, geodetic, and tsunami observations. Here we show that among six tested source models, there is a mean 47% gain in positively stressed aftershock mechanisms over that for the background (1997-10 March 2011) earthquakes, which serve as the control group. An aftershock fault friction of 0.4 is found to fit the data better than 0.0 or 0.8, and among all the tested models, Wei and Sladen (2011) produced the largest gain, 63%. We also calculate that at least 5 of the seven large, exotic, or remote aftershocks were brought ≥0.3 bars closer to failure. With these tests as confirmation, we calculate that large sections of the Japan trench megathrust, outer rise normal faults, the Kanto fragment beneath Tokyo, and the Itoigawa-Shizuoka Tectonic Line, were also brought ≥0.3 bars closer to failure.

1. Introduction

The M=9.0 Tohoku-chiho Taiheiyo-oki (hereafter, 'Tohoku') earthquake is unprecedented in size in Japan's long recorded history of earthquakes, although some foresaw its possibility (Kanamori *et al*., 2006; McCaffrey, 2008). The earthquake struck

on the Japan trench megathrust, which accommodates ~80 mm/yr of convergence. Thanks to Japan's superb monitoring networks, the earthquake is the best-recorded great event the world has ever known. Seismic, geodetic and tsunami observations not only permit detailed source inversions for the distribution of slip on the megathrust surface, but also provide an unparalleled set of aftershocks, which the NIED has used to calculate focal mechanisms for M≥3.5 events recorded by F-net. Further, the extensive geomorphic, geodetic and paleoseismic analysis of Japan's faults enables us to calculate how the Tohoku earthquake changed the conditions for failure on surrounding faults.

2. The Coulomb Stress Triggering Hypothesis

An earthquake fault rupture permanently deforms the surrounding crust, changing the stress on nearby faults as a function of their location, geometry and sense of slip (rake). The Coulomb stress change is defined as $\Delta CFF = \Delta\tau + \mu\Delta\sigma$, where τ is the shear stress on the fault (positive in the inferred direction of slip), σ is the normal stress (positive for fault unclamping), and μ is the apparent friction coefficient. Failure is promoted if ΔCFF is positive, and inhibited if negative; both increased shear and unclamping of faults are taken to promote failure, with the role of unclamping modulated by fault friction. Most investigations of Coulomb stress triggering (Harris, 1998; Stein, 1999; Freed, 2005) find that static stress change plays an important role in the production of aftershocks and subsequent mainshocks on surrounding faults.

3. Test Design

While there is abundant evidence that aftershocks locate in regions of calculated stress increase (the stress triggering lobes) for assumed 'receiver fault' geometries, a stricter test of the Coulomb hypothesis is whether the nodal planes of the aftershocks are promoted for failure. By resolving the Coulomb stress change on aftershock nodal planes in their rake directions, one need not make any assumptions about the aftershock fault geometry or the regional stress. But while the shear stress on the two orthogonal nodal planes is the same, the unclamping stress is different. So, except for the special case of zero fault friction, the Coulomb stress imparted to the two nodal planes will differ, and except under unusual circumstances we do not know which of the two planes slipped. Thus to conduct this test, we resolve the Coulomb stress change on both nodal planes of each mechanism, and use both resulting stress changes in the statistical sample. This

means that because of the nodal plane ambiguity, even if all aftershocks were brought closer to failure we would never find a 100% accord.

But even if we found that a majority of aftershock nodal planes were brought closer to failure by the mainshock rupture, this would not necessarily demonstrate that stress transfer was responsible. It is possible that shocks that occurred before the mainshock rupture in the same region would yield a similar percentage, principally because faults that slip during aftershocks might share the same geometry as those that slip at other times. Therefore, the percentage of positively stressed aftershocks must be normalized by the positively stressed background shocks. For the 'background,' we use shocks that occurred before the mainshock. In testing terminology, the aftershocks comprise the 'test group,' and the background shocks the 'control group.' What is important is the percentage gain in positively stressed aftershocks relative to the background shocks, with the background shocks selected to match the geographic area and depth range as the aftershocks. This procedure, first used by Hardebeck *et al*. (1998), is analogous to pharmaceutical testing, in which to prove efficacy, the drug given to the test group must outperform a placebo administered to the control group.

It has proven difficult to discriminate among possible values of fault friction in Coulomb stress transfer studies, and so a mid-value of 0.4 is most commonly adopted. On creeping faults, low friction has been found to fit best, whereas on young thrust and normal faults, a high value of friction may be superior (Parsons *et al*., 1999; Toda and Stein, 2002). Because of the quality of the test data for the Tohoku earthquake, we test three values of fault friction (0.0, 0.4, and 0.8) for each source model, seeking the value that produces the highest gain.

To graphically represent the results economically, we plot the most positive stress change of each pair of nodal planes in Fig. 1, with the convention that where aftershocks overlap, the most positively-stressed shocks are plotted on top. This schema is used for both the aftershocks (Fig. 1a) and the background shocks (Fig. 1b); this 'positive bias' is only in the figure; stress on both planes are used in all calculations, as summarized in Table 1. We performed all calculations using Coulomb 3.2 (http://www.coulombstress.org).

4. Aftershock Data and Source Models

We use the 177 NIED F-net aftershock focal mechanisms between 11 and 31 March 2011; all are M≥3.5 events within the depth range of 0-80 km, bounded by 34.5-41.0°N latitude and 137-145°E longitude (Fig. 1a). For the background earthquakes, we restricted our search area and depths to the same range as the aftershocks, and sought a dataset about five times larger than the aftershocks for statistical confidence; we used 840 M≥4.5 NIED F-net focal mechanisms during 17 February 1997-10 March 2011 (Fig. 1b). We chose M≥4.5 to limit the size of the sample, and because larger earthquakes likely have more accurate focal mechanisms.

We tested six source models of the 11 March 2011 mainshock that span the range of the assumed megathrust dips (9-15°) and datasets used in the source inversions (teleseismic P, SH, body, and long period surface waves; static GPS displacements; and tsunamigrams and coastal tide gauge records), as noted in Table 1. First, we use the combined statistics for all models to test the Coulomb hypothesis, and to find the best value of fault friction. We then use the test to discriminate among the candidate models, with the expectation that the source model producing the highest gain in positively stressed aftershocks relative to background shocks will provide the best estimates of which major surrounding faults have been brought closer to failure by the Tohoku earthquake.

5. Test Results and their Implications

The mean gain in the percentage of positively-stressed aftershocks with respect to the background shocks for all six source models is 47% (47% more aftershocks are brought closer to failure than the control group). We find a friction of 0.4 yields a 5-10% higher gain than for low or high friction (Table 1). Since some aftershocks locate on or near the megathrust surface and others locate in the outer rise or on continental faults (Fig. 1a), the 0.4 preference could simply be due to averaging different fault strengths. Among the candidate source models, Wei and Sladen (2011) (Figs. 1 and 2) has the highest mean gain for all values of friction (55%); it has the highest gain of all (63%), obtained for friction=0.4. We thus adopt this model and friction for the remaining tests, and for calculations of stress transfer to surrounding faults. Although many of the aftershocks are on the megathrust, only 6% of the stress changes exceed |50 bar|, and so proximity to the rupture surface does not strongly influence the results.

In a similar test, Hardebeck *et al*. (1998) found a 37% gain for immediate aftershocks of the 1992 M=7.3 Landers, California, earthquake, and 46% gain after 3 years. Ma *et al*. (2005) found a gain of 61% (70% for thrust and a 56% gain for strike-slip events) for the 1999 M=7.6 Chi-Chi, Taiwan, earthquake, using a 4-yr-long aftershock period. These studies tested only a single source model, and were unable to discriminate among friction values. Because they lack a sufficient number of pre-mainshock focal mechanisms, such tests have not been conducted for the 2004 M=9.1 Sumatra, 2008 M=7.9 Wenchuan, and 2010 M=8.8 Chile earthquakes.

We next examine the Coulomb stress triggering to the largest, most exotic, and remote aftershocks. For this test we use the Global CMT catalog, which is more complete than the NIED F-net catalog during the first day after the mainshock. The events include the Mw=7.6 outer rise aftershock, several shallow coastal and offshore 5.4≤Mw≤5.8 normal events, the thrust event near Nagano, and the remote strike-slip events in the Japan Sea and at the base of Mt. Fuji (Fig. 2). We find that all but the Nagano event (aftershock #2 in Fig. 2 and Table 2) are brought ≥0.3 bars closer to failure on at least one nodal plane, and that four of the 7 are brought closer to failure on both nodal planes. Thus, the Nagano event does not appear to be statically triggered. 0.3 bar is about 0.25% of the Tohoku earthquake stress drop of ~120 bar for the Wei and Sladen (2011) model, and about three times larger than the minimum stress typically found to cause changes in seismicity rates (Stein, 1999).

6. Stress Changes Calculated on Major Faults in Central Japan

Emboldened by the success of the Coulomb hypothesis test for aftershocks, we next calculate the Coulomb stress changes on all known major faults in Fig. 3 to develop insights as to which fault systems are now more hazardous than they were before the Tohoku earthquake. We nevertheless caution that the calculation idealizes stress transfer because smaller faults need not share the same geometry as the major faults sampled here, and the dip and rake of even major faults are often poorly known.

On the megathrust surface (which is rather unrealistically represented as a plane in all Tohoku source models), there are large stress increases at depths greater than 35 km, which could give rise to aftershocks and postseismic slip. To the north of the Tohoku rupture in the Sanriku-Hokubu-oki area, there are also large stress increases at depths

greater than 30-35 km (Fig. 3). This section hosted the 1994 Mj=7.5, 1901 Mj=7.4, 1931 Mj=7.6, and 1968 Mj=7.9 and 7.5 earthquakes (south to north respectively). To the south of the Tohoku rupture lies the Off-Boso section (Fig. 3), which exhibits repeating earthquakes, aseismic slip transients, and possibly uncoupled behavior (Nishimura *et al*., 2006; Ozawa *et al*., 2007). Despite this, there have been 2 M>6 and 4-5 M≥5 Off-Boso aftershocks of the Tohoku earthquake (Fig. 2), and so it must also store elastic stress. Stress increases of about 10 bars are calculated along the northern margin of the Off-Boso section of the Japan trench adjacent to the M=7.9 aftershock, declining to 0.3 bars to the south. If this entire 100- by 180-km section ruptured, it could host a M=8.1 earthquake, and so we believe that such a possibility cannot be dismissed even though no such event is evident in the historical record (Grunewald and Stein, 2006).

The sections of the Sagami trough that last ruptured in the 1923 M=7.9 Kanto event (Nyst *et al*., 2006) are calculated to experience Coulomb stress decreases of 0.2-0.5 bar by the Tohoku earthquake, whereas the stress increased by 0.5-1.0 bar on the easternmost section. The stressed sections might have participated in the 31 December 1703 M~8.2 Genroku and 4 November 1677 M~8? tsunamigenic earthquakes (Grunewald and Stein, 2006). The Suruga trough megathrust (Ando, 1975; Ishibashi, 1981), site of the 1854 M=8.4 event, was brought a negligible 0.02-0.07 bars closer to failure since it is not much larger than the 0.01-bar Coulomb stresses induced by the tides. The left-lateral Itoigawa-Shizuoka Tectonic Line (ISTL), as well as many parallel faults to the west of it, were brought 0.3-0.4 bar closer to failure. For a fault friction of 0.4, the major thrust faults of Tohoku were inhibited from failure by 1 bar. However, for a high value of friction that might be more appropriate for youthful thrust continental thrusts (Parsons *et al*., 1999), some Tohoku thrusts are unclamped, leading to Coulomb stress increases of 0.1-0.4 bar (see Fig. 3 inset). In contrast to the Tohoku thrusts, failure is promoted on outer rise normal faults by 1.5-15 bars, depending on their proximity to the locus of high slip. Numerous large Tohoku outer rise normal aftershocks have struck here since the Tohoku mainshock, including the 11 March 2011 Mw=7.6 event (Fig. 2). The largest outer rise event in this region was the 3 March 1933 M=8.1 quake, whose tsunami inundated the Sanriku coast (Kanamori, 1971). The southern 2/3 of the likely 1933 rupture zone (long red rectangle near the trench in Fig. 3) was brought 6 bars closer to failure.

The ‘Kanto fragment’ was proposed by Toda *et al*. (2008) to explain the seismic tomography, microseismicity, geodetic deformation and plate motion evolution of the

Philippine Sea and Pacific slab interaction beneath Tokyo. They argued that the fragment broke off the Pacific slab and is wedged between the Pacific and Philippine Sea slabs and the over-riding Eurasian plate. We calculate that the 60-80-km-deep base of the fragment has been brought 1-2 bars closer to failure, whereas the upper surface, in contact with the Philippine Sea slab, has been brought 0.3 bars closer to failure. If the entire lower surface of the fragment were to rupture, a M=7.3 event could strike at ~75 km depth beneath the highly-populated Kanto basin.

7. Conclusions

We have strived to conduct the most rigorous test possible of the Coulomb failure hypothesis that can be applied to the Tohoku earthquake less than one month after the mainshock. The test is made feasible by the extraordinary quality and accessibility of the Japanese seismic and geodetic monitoring networks, and by scientists openly sharing their preliminary source models, for which we are grateful. We find that all source models yield a at least a 43% gain in positively stressed aftershock nodal planes relative to the nodal planes of background seismicity, and the best tested model yields a 63% gain, the highest yet seen for any earthquake tested. When this model is used to examine the seven remote, exotic, and largest aftershocks, at least one nodal plane of six of these is brought >0.3 bar closer to failure, and so the remote aftershocks can have the same static-stress origin as those on the rupture surface. Examining all major fault systems in central Japan, large stress increases are calculated to the north and south and along the base of the Tohoku megathrust rupture surface, on outer rise normal faults, on the easternmost sections of the Sagami trough megathrust, along the ISTL and sub-parallel faults, and on the base of the Kanto fragment that, we believe, underlies Tokyo.

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Figure captions

Figure 1. Stress imparted by the Wei and Sladen (2011) source model and the M=7.9 aftershock are resolved on both nodal planes of NIED F-net aftershock (**a**) and background, or pre-Tohoku (**b**) focal mechanisms. Although both planes are used for the calculations summarized in Table 1, the most positively-stressed of each pair of nodal planes is shown here. The effect of the mainshock stress imparted to the mechanisms is judged from the percent gain in positively stressed mechanisms after the mainshock (the test group) relative to the background shocks (the control group). Results for all 6 models are given in Table 1.

Figure 2. Focal mechanisms of the March 2011 Tohoku earthquake aftershocks from two catalogs, with important aftershocks examined for stress transfer keyed to Table 2. The mainshock promotes failure of five of the seven events regardless of nodal plane or friction; for the remaining two events (#1-2) only one of the nodal planes is promoted (see Table 2). All aftershocks east of the mainshock are outer rise normal events, but there is also a group of coastal normal events midway between Tokyo and Sendai.

Figure 3. Stress imparted by the Wei and Sladen (2011) source model and the M=7.9 aftershock to surrounding active faults (Research Group for Active Faults in Japan, 1991; Headquarters for Earthquake Research Promotion, 2011), resolved in their inferred rake directions (oblique rakes are labeled). Top depth and bottom depths of most of the active faults are set to 0 and 15 km.

Table Captions

Table 1. Coulomb stress triggering test results. A representative sample of six inversion models were tested on NIED F-net aftershocks and background shocks. The background shocks were chosen to sample the same depth range and area as the Tohoku aftershocks. All models showed gains of 43-55% over the control (right-hand column). Except if friction=0.0, it is impossible to have positive stress change on 100% of nodal planes, because the Coulomb stress change to each nodal plane in a pair is different, and we do not know which is the fault plane. The best value of fault friction among all models is about 0.4, and the best model among all models is Wei and Sladen (2011).

Table 2. Coulomb stress imparted to the nodal planes of large, exotic, or more remote aftershocks of the Tohoku earthquake. The earthquake number is keyed to Fig. 2. All but aftershock #2 is brought ≥0.3 bars closer to failure on at least one nodal plane.

Table 1

Source model [1] (2011 citation)	**Ver.**	**Moment** E29 dyn-cm	**Dip** (°)	**Slip patches**	**Data** [2]	**% aftershocks with +ΔCFF** [3] fric.=0.0	fric.=0.4	fric.=0.8	**% background with +ΔCFF** [4] fric.=0.0	fric.=0.4	fric.=0.8	**Stress effect of mainshock (% gain)** [5] fric.=0.0	fric.=0.4	fric.=0.8	**Model means**
Yagi	v.2	4.5	15	264	a	63	67	69	45	45.8	48.6	56.6	52.9	40.6	50.0
Hayes	v.2	4.9	10.2	325	b	74	76	73	51.5	49.8	51.8	40.0	46.3	42.0	42.8
Shao et al	v.3	5.7	10	190	b	64	63	61	45.5	41.2	43.4	43.7	52.6	40.9	45.7
Fuji & Satake	v.1	2.8	14	25	c	77	81	82	55.7	55.8	56.6	38.2	45.2	44.9	42.8
Wei & Sladen	v.1	4.5	9	350	b, d	73	74	74	50.2	45.5	47.5	45.4	**62.6**	55.8	54.6
Pollitz & Burgmann	v.1	3.6	14	446	d, e	68	67	68	42.1	43.9	51.9	61.5	52.6	31.0	48.4
											Friction means:	47.6	52.0	42.5	47.4

Notes:

[1] M=7.9 aftershock included as a source (uniformly tapered with 5 patches from this study, based on Global CMT parameters and ARIA GPS; Mo=8.5E27 dyn-cm)

[2] a, teleseismic body waves; b, teleseismic P, SH, and long period surface waves; c, DART tsunamigrams and tide gauges; d, static GPS; e, smoothed from 15,876 patches

[3] 176 aftershocks used from NIED F-net database for 3/11/2011-3/31/2011 (all are M≥3.5, 0-80 km depth, and lie within lat. 34.5-41.0°N, lon. 137-145°E

[4] 840 background shocks from NIED F-net database for 2/17/1997-3/10/2011 (selection criteria: M≥4.5, 0-80 km depth, lat. 34.5-41.0°N, lon. 137-145°E)

[5] Stress effect of mainshock=[100*(% aftershocks positively stressed by mainshock/% background mechanisms positively stressed by mainshock)-100]; so if 0%, there is no effect of the rupture on aftershock mechanisms, and if 100%, aftershock mechanisms are promoted at twice the background rate.

Table 2

#	Occurrence time (local time) mo/dy/yr hr:mn	Lon. (°)	Lat. (°)	Depth (km)	Mw	Mj	strike NP1 (°)	dip NP1 (°)	rake NP1 (°)	normal stress (bar)	shear stress (bar)	Coulomb stress μ=0.4 (bar)	Coulomb stress μ=0.8 (bar)	strike NP2 (°)	dip NP2 (°)	rake NP2 (°)	normal stress (bar)	shear stress (bar)	Coulomb stress μ=0.4 (bar)	Coulomb stress μ=0.8 (bar)
1	3/11/11 15:26	144.63	38.27	21.1	7.6	7.5	182	42	-100	**-9.3**	**2.2**	**-1.5**	**-5.2**	15	49	-81	**4.9**	**2.2**	**4.1**	**6.1**
2	3/12/11 3:59	138.59	37.08	12.0	6.3	6.7	28	33	58	**0.7**	**-0.3**	**0.0**	**0.3**	244	63	109	**0.1**	**-0.3**	**-0.3**	**-0.2**
3	3/12/11 4:46	139.15	40.4	12.0	6.2	6.4	27	78	-177	**1.9**	**0.5**	**1.2**	**2.0**	296	87	-12	**0.0**	**0.5**	**0.5**	**0.5**
4	3/15/11 22:31	138.65	35.29	17.7	5.9	6.4	296	70	172	**0.7**	**0.4**	**0.6**	**0.9**	29	82	20	**-0.1**	**0.4**	**0.3**	**0.3**
5	3/19/11 18:56	140.55	36.85	12.0	5.8	6.1	146	44	-81	**3.2**	**3.9**	**5.1**	**6.4**	314	47	-98	**3.3**	**3.9**	**5.2**	**6.5**
6	3/23/11 7:12	140.78	37.09	12.0	5.7	6.0	15	37	-92	**5.4**	**6.6**	**8.8**	**11.0**	197	53	-89	**9.0**	**6.6**	**10.2**	**13.8**
7	3/23/11 7:36	140.76	37.05	12.4	5.4	5.8	182	42	-92	**6.4**	**7.1**	**9.7**	**12.3**	5	48	-88	**8.3**	**7.1**	**10.4**	**13.8**

Wei & Sladen (2011) plus the Mw=7.9 aftershock model from this study are used as sources. NP1 = Nodal Plane 1, etc. μ = friction coefficient.

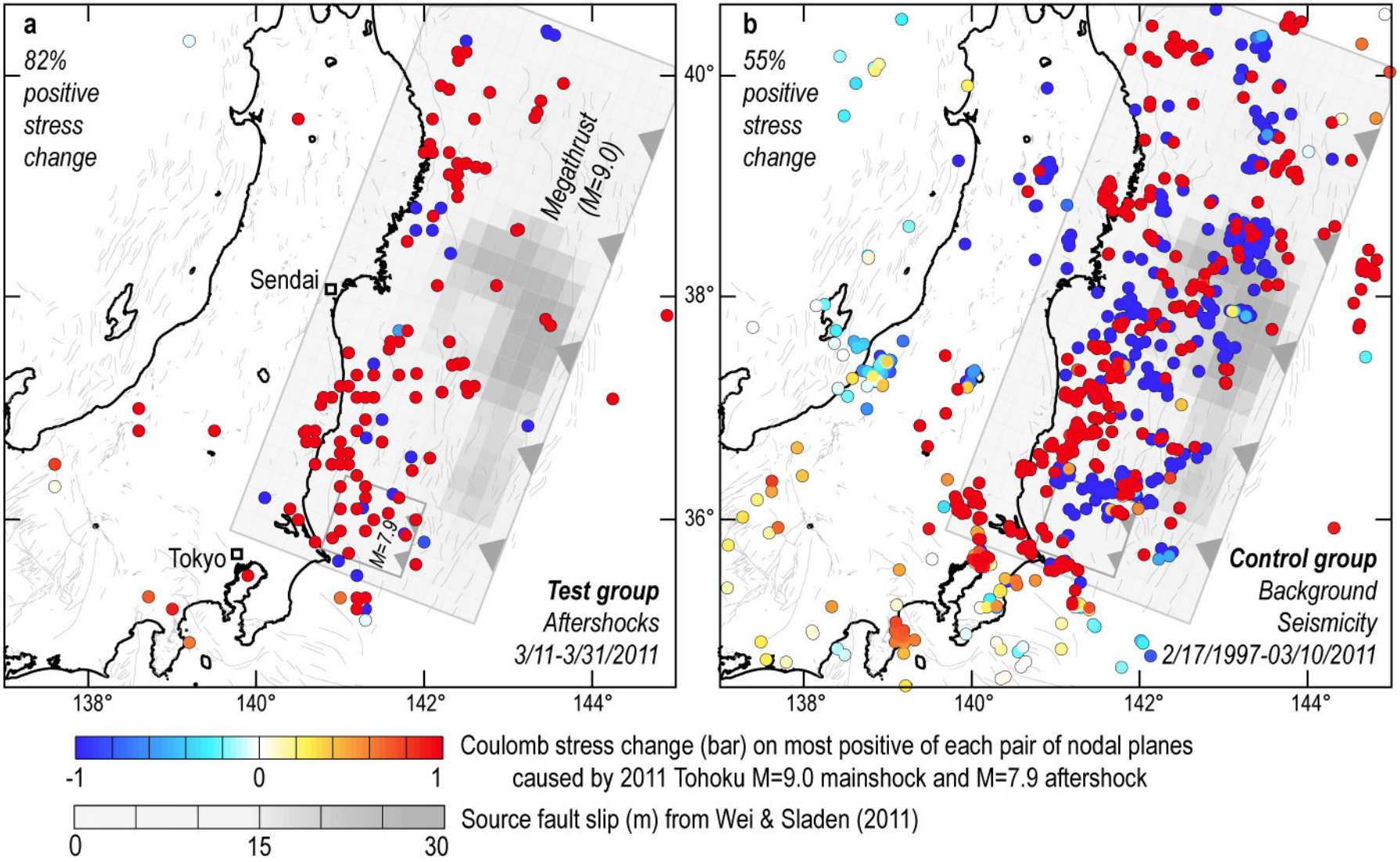
a
82%
positive
stress
change
Megathrust
(M=9.0)
Sendai
Tokyo
M=7.9
Test group
Aftershocks
3/11-3/31/2011
b
55%
positive
stress
change
Control group
Background
Seismicity
2/17/1997-03/10/2011
40°
38°
36°
138°
140°
142°
144°
-1
0
1
Coulomb stress change (bar) on most positive of each pair of nodal planes
caused by 2011 Tohoku M=9.0 mainshock and M=7.9 aftershock
0
15
30
Source fault slip (m) from Wei & Sladen (2011)

Fig. 1

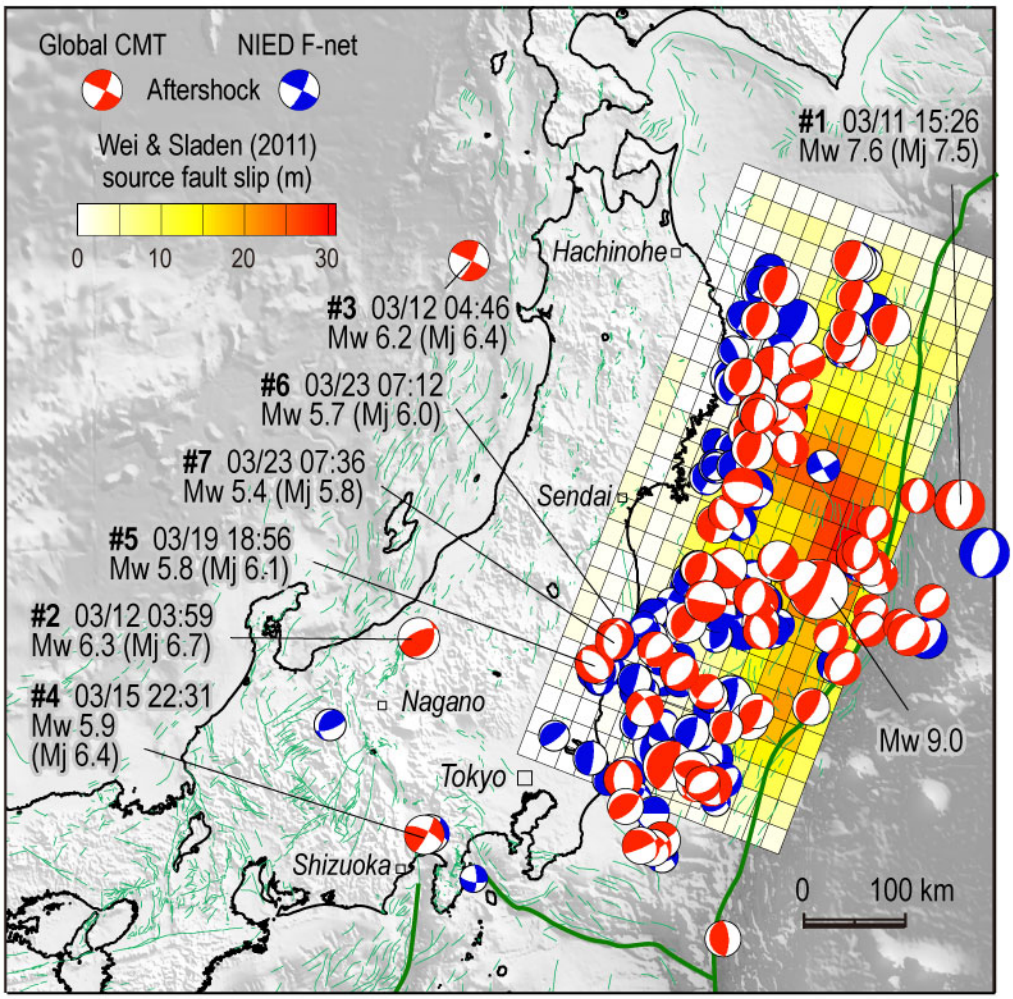
Global CMT
NIED F-net
Aftershock
Wei & Sladen (2011)
source fault slip (m)
0
10
20
30
#1 03/11 15:26
Mw 7.6 (Mj 7.5)
Hachinohe
#3 03/12 04:46
Mw 6.2 (Mj 6.4)
#6 03/23 07:12
Mw 5.7 (Mj 6.0)
#7 03/23 07:36
Mw 5.4 (Mj 5.8)
Sendai
#5 03/19 18:56
Mw 5.8 (Mj 6.1)
#2 03/12 03:59
Mw 6.3 (Mj 6.7)
#4 03/15 22:31
Mw 5.9
(Mj 6.4)
Nagano
Tokyo
Shizuoka
Mw 9.0
0
100 km

Fig. 2

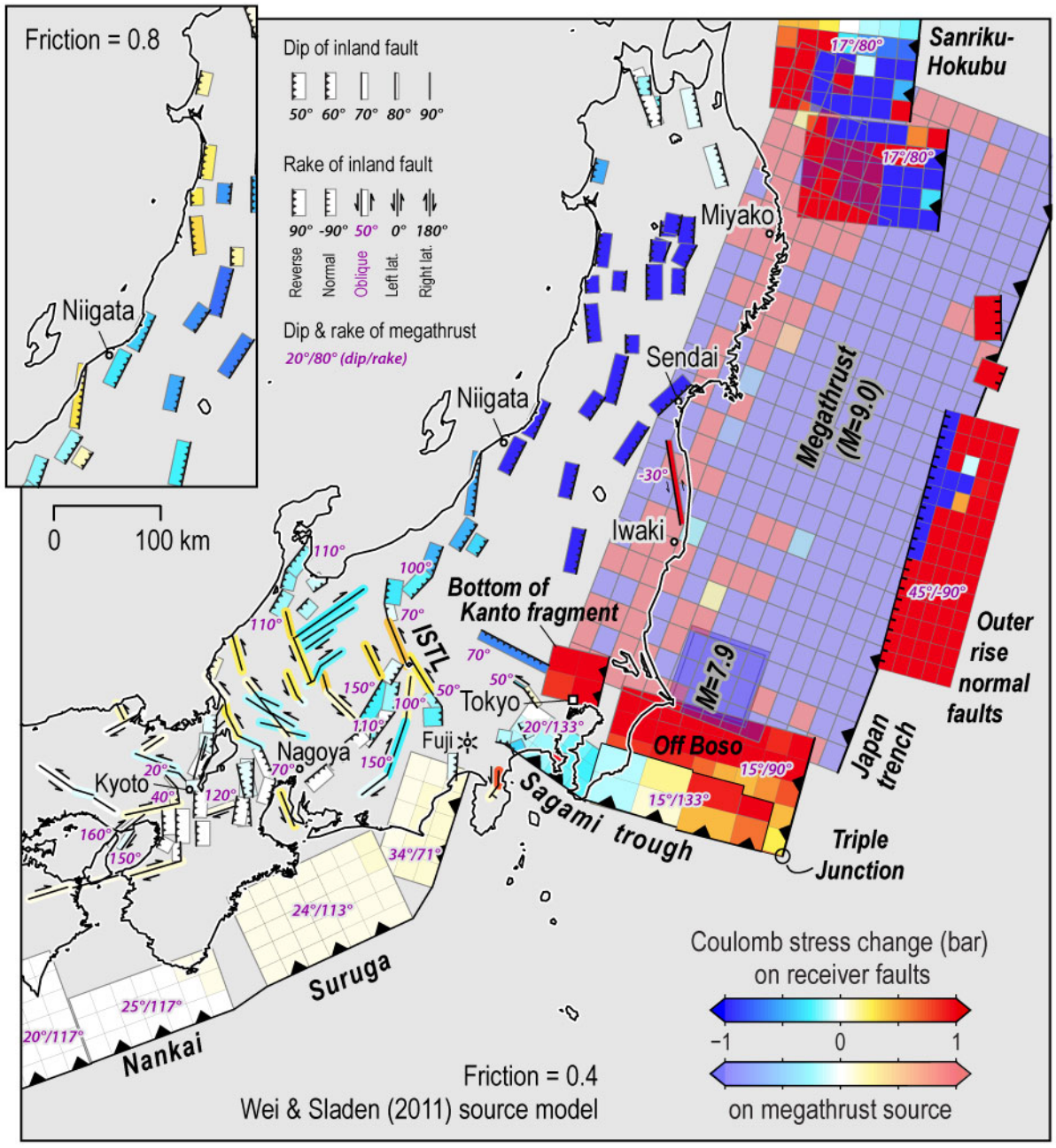
Friction = 0.8
Niigata
0
100 km
Dip of inland fault
50° 60° 70° 80° 90°
Rake of inland fault
90° -90° 50° 0° 180°
Reverse
Normal
Oblique
Left lat.
Right lat.
Dip & rake of megathrust
20°/80° (dip/rake)
Niigata
Miyako
Sendai
Iwaki
Tokyo
Fuji
Nagoya
Kyoto
Sanriku-Hokubu
17°/80°
17°/80°
Megathrust (M=9.0)
-30°
45°/-90°
Outer rise normal faults
Bottom of Kanto fragment
M=7.9
Off Boso
15°/90°
15°/133°
Japan trench
Triple Junction
Sagami trough
ISTL
20°/133°
34°/71°
24°/113°
25°/117°
20°/117°
Suruga
Nankai
110°
100°
70°
50°
150°
20°
40°
120°
160°
Coulomb stress change (bar) on receiver faults
−1 0 1
on megathrust source
Friction = 0.4
Wei & Sladen (2011) source model

Fig. 3